

# Basic Glaciology

Winter School, February 2009

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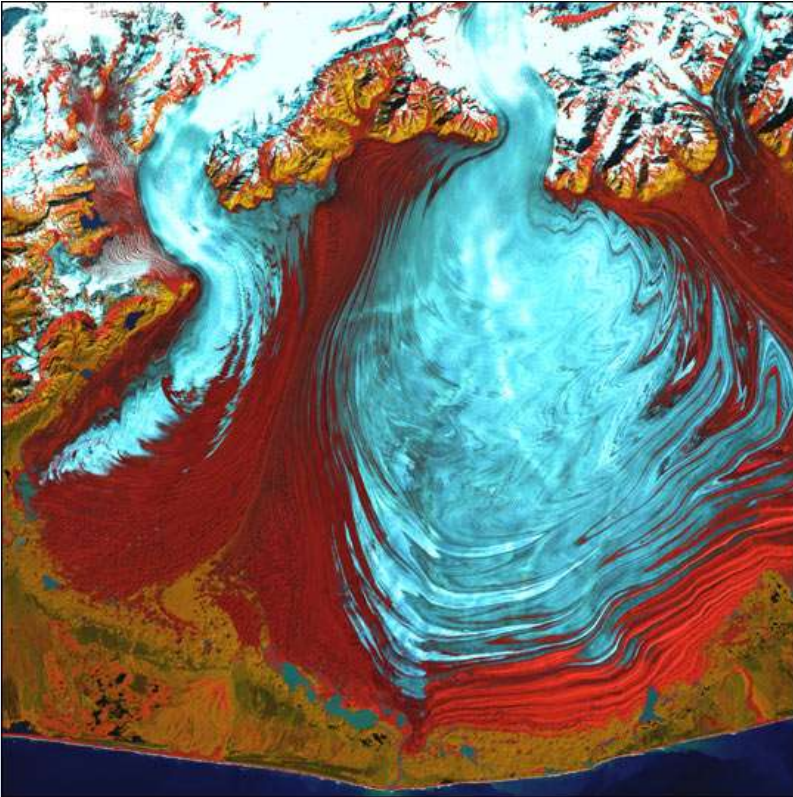


# Outline

- What is a glacier?
- Formation
- Flow
- Advance and retreat
- Erosion
- Glacial geomorphology

# What is a glacier:

Natural body of ice that flows under its own weight.



LANDSAT image of Malaspina Glacier AK, NASA visible earth



Glikey Glacier, AK, Mellisa Rhody

Observed thicknesses:  $>4\text{km}$  to  $\sim 30\text{m}$

Observed surface velocities:  $<1\text{mm}$  to  $>10\text{m}$  per day

# Formation: accumulation and ablation

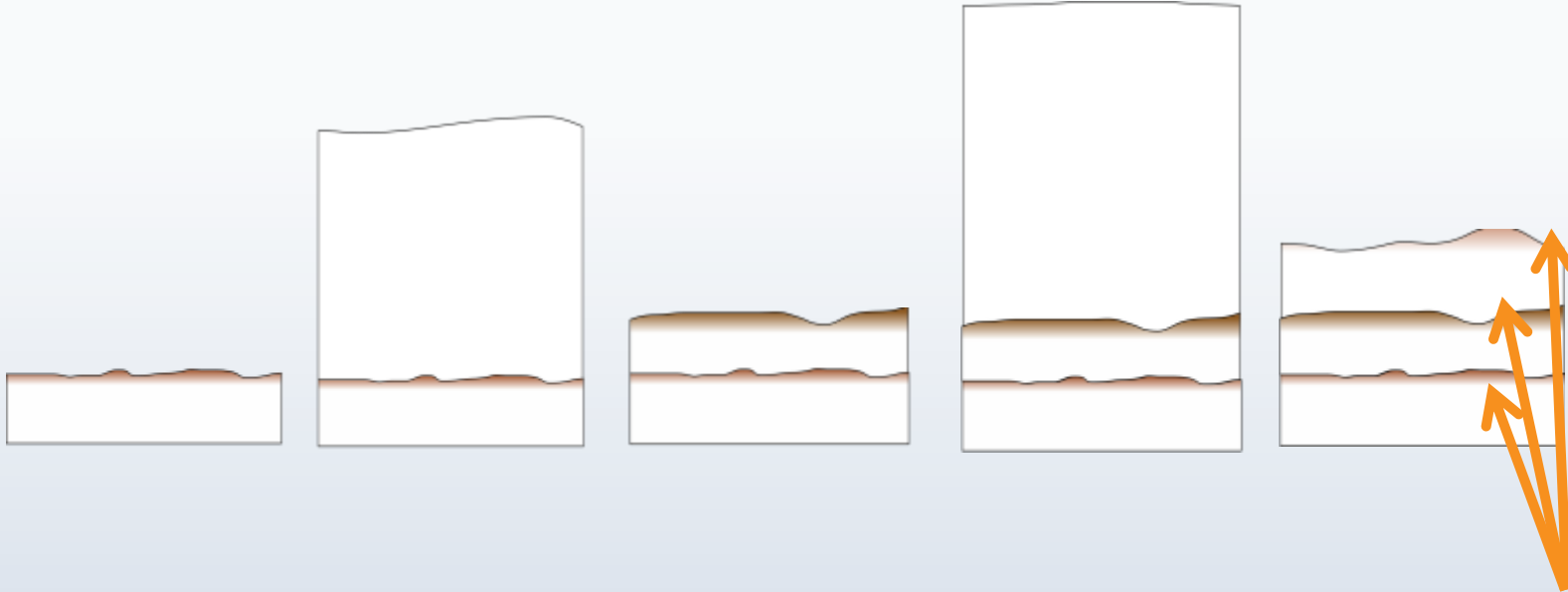
Year 1:  
Summer

Year 1:  
Winter

Year 2:  
Summer

Year 2:  
Winter

Year 3:  
Summer



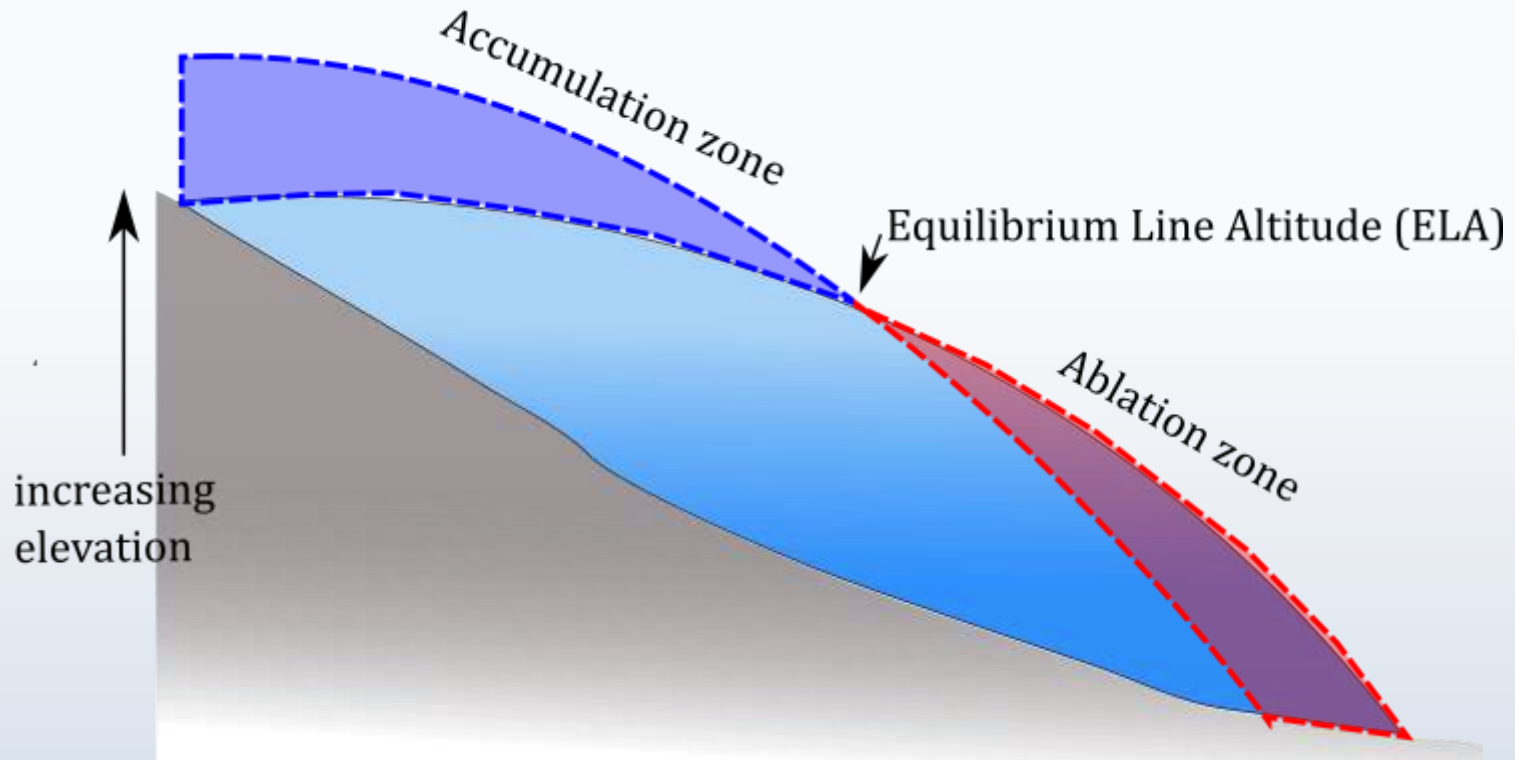
Glacial ice is formed by a net accumulation of snow

"annual layers"

The natural timestep for glacial systems is 1 year

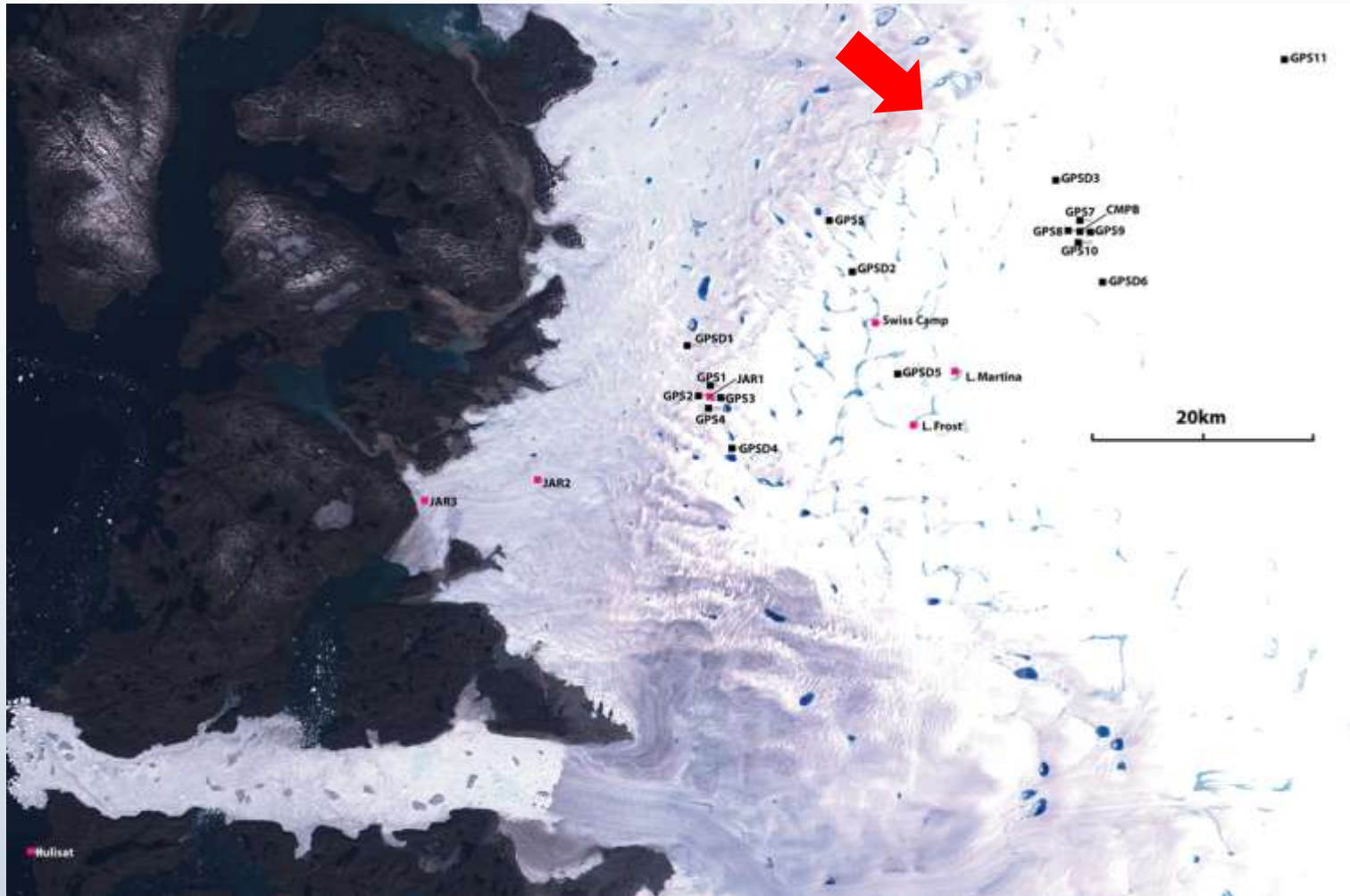
This snow becomes ice by mechanical breakdown, melting and refreezing, and recrystallization under pressure

# Formation: mass balance



- Temperature and precipitation vary with elevation, so does mass balance
- 3 features: Accumulation zone, ablation zone, and ELA

# Glacier formation: examples



Jacobshavn Glacier, MODIS, [http://www.ig.utexas.edu/research/projects/meltwater\\_greenland/](http://www.ig.utexas.edu/research/projects/meltwater_greenland/) Aerial photo by J. Alton, SwissEduc

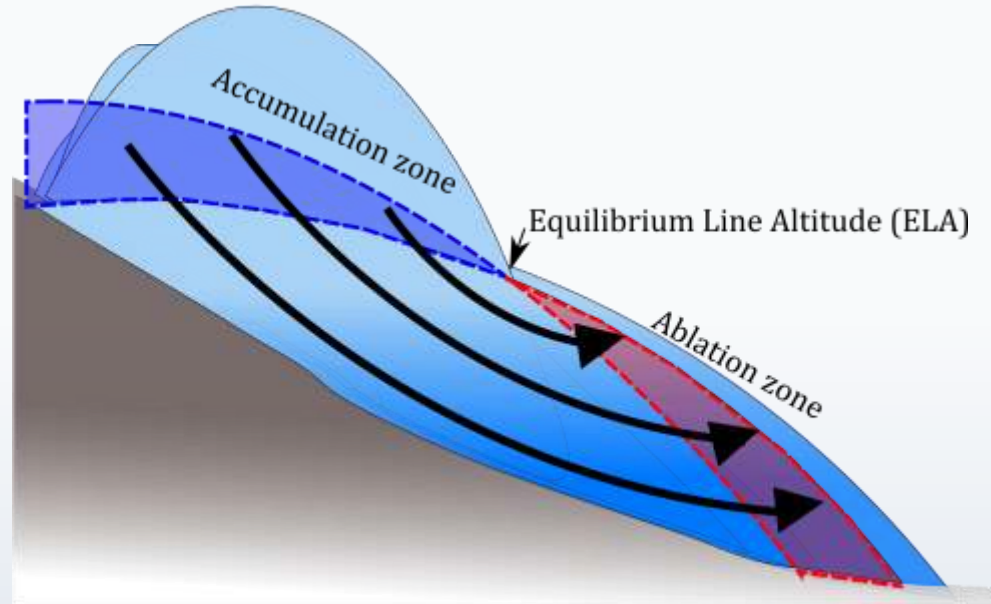
- These zones apply to glaciers of all sizes, and can be easily seen

# Measuring net accumulation



- Digging snow pits to find the annual layer and measure the density of the firn

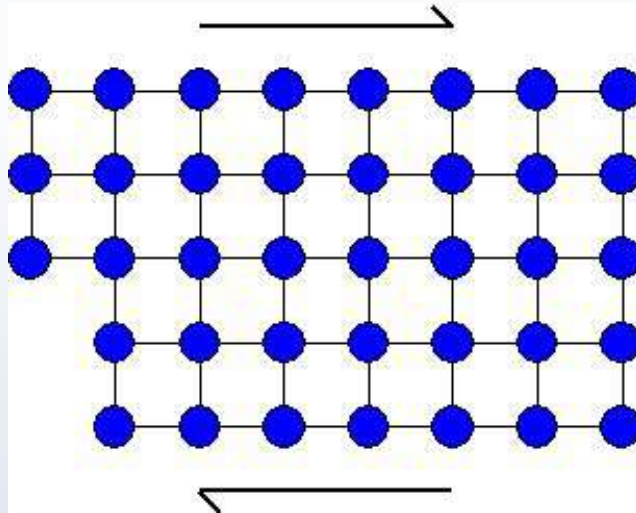
# Mass balance and equilibrium form



- Static ice would not sustain an ablation zone...
- Ice flow moves mass from accumulation to ablation zone, controlling glacier shape and volume
- If climate forcing is constant, the glacier will reach and equilibrium form

# Flow: ice deformation

- Ice deforms by dislocation creep



- A simple empirical relation describes this type of deformation:

# Flow: ice deformation

- $\tau$  is stress (force/area) parallel to the ice surface

Patterson figure here

- $A$  is a rate constant, defined by the an Arrhenius relation:

# Flow: ice deformation

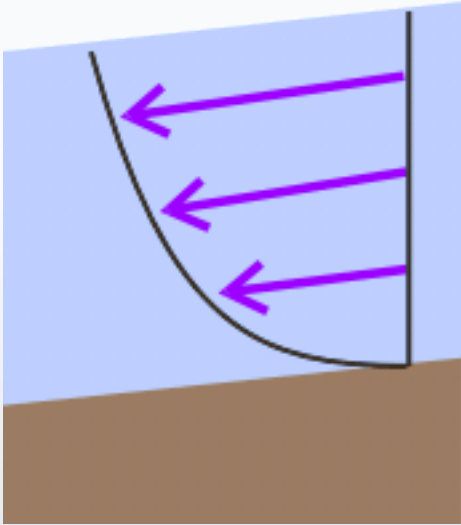
- Substituting gives:

→ velocity depends on temperature (T), ice thickness (h) and surface slope ( $\alpha$ ), basically on mass balance and climate!

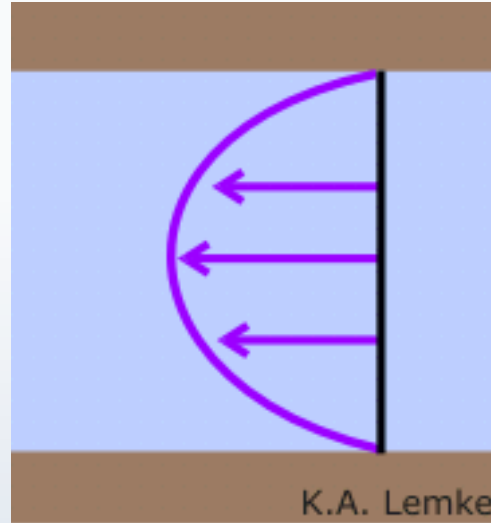
- Cold ice flows *much* more slowly
  - 10°C : 10x slower than 0°C
  - 25°C : 100x slower than 0°C

# Flow: velocities

Vertical Profile



Transverse Profile

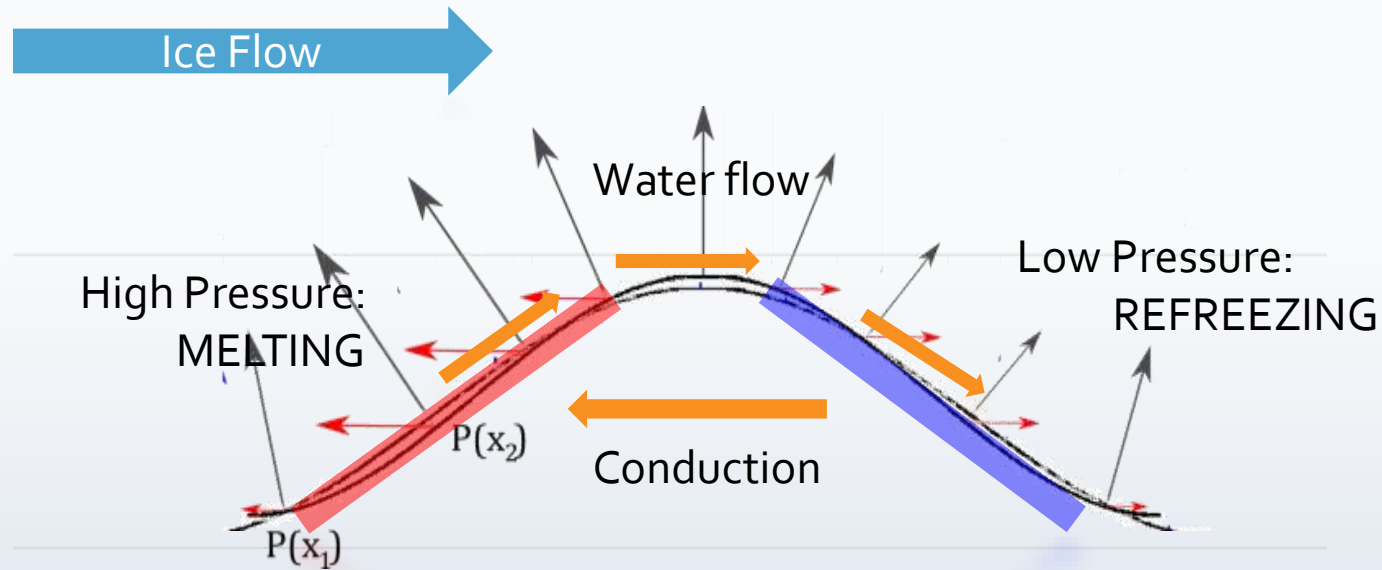


K.A. Lemke

Deformation only case:

- Most deformation occurs at the base and sides
- Highest velocities occur at the surface and centerline

# Flow: Basal Sliding

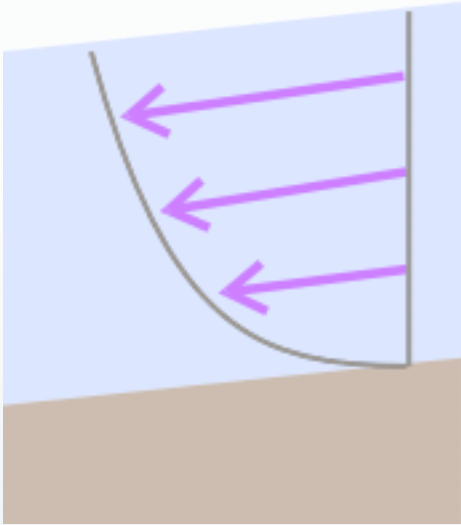


Schematic of stresses on ice components in the direction of ice flow

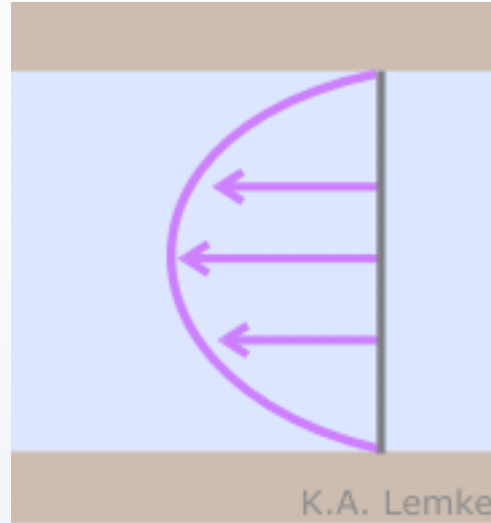
- Bedrock "obstructions" resist sliding
- Ice is able to move past by regelation: pressure melting and refreezing
- This process depends on temperature and  $\tau$

# Flow: velocities

Vertical Profile

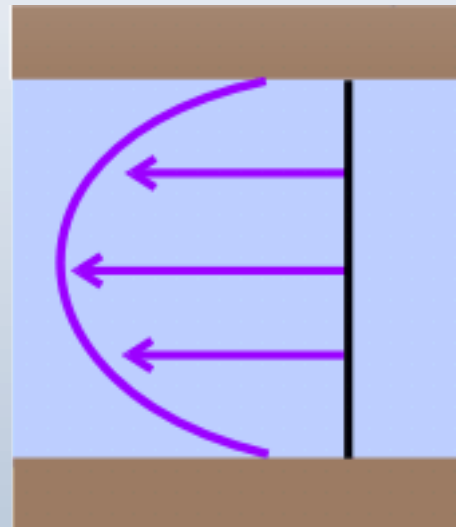
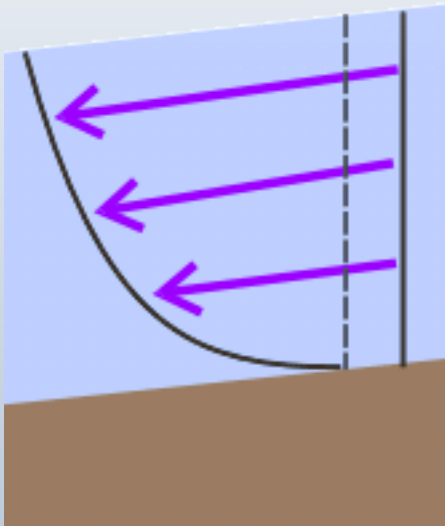


Transverse Profile



Deformation only case:

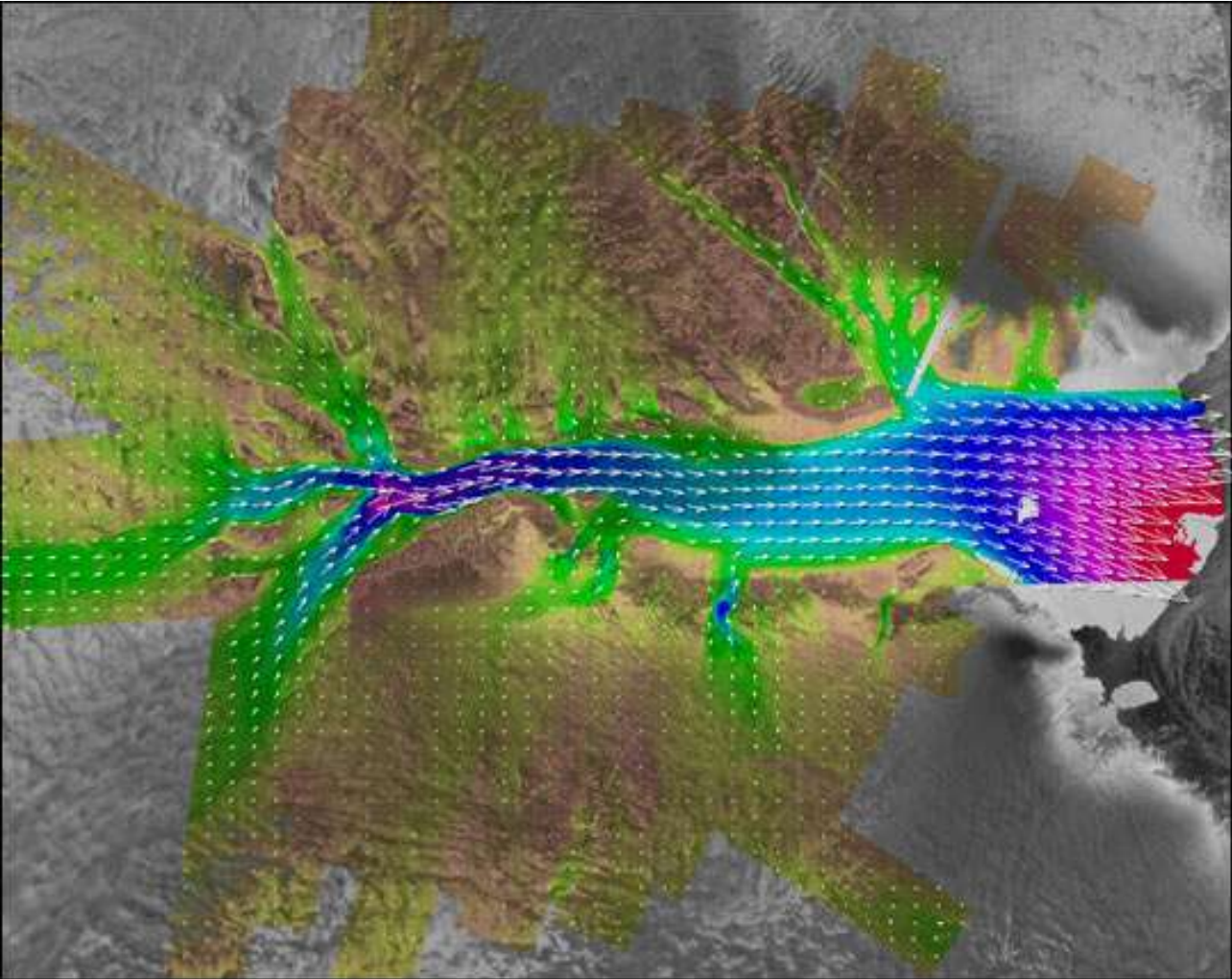
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Including sliding:

- Non-zero velocity at base and at valley walls

# Flow: velocities

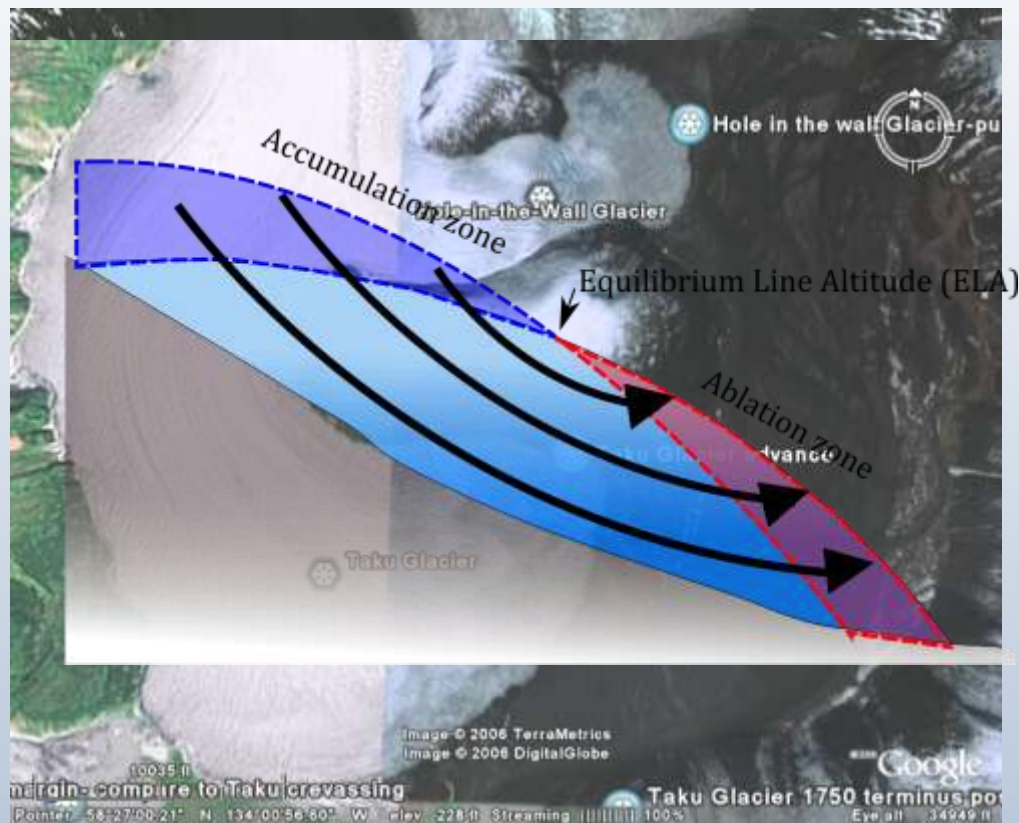


- Some velocity data

GPS velocity measurements of ice velocities, NSIDC

# Advance and retreat

- Mass balance ultimately drives extent
- Glacier extent records a combination of temperature and precipitation
- Each glacier responds to different forcing



# Mechanisms of erosion:

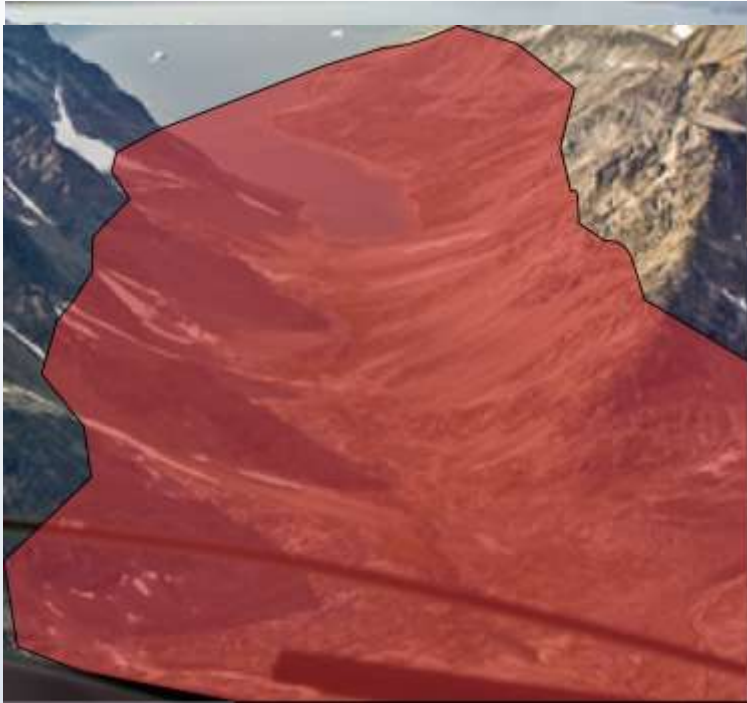
- Abrasion
- Crushing
- Quarrying
- Subglacial fluvial erosion
- Chemical erosion (dissolution)



New Zealand, Tom Lowell's Glacier Image Database, University of Cincinnati

Mt Lyell, Paul Doherty 2003

# Glacial geomorphology: valley shape



Greenland, "Olikristinn", Flickr

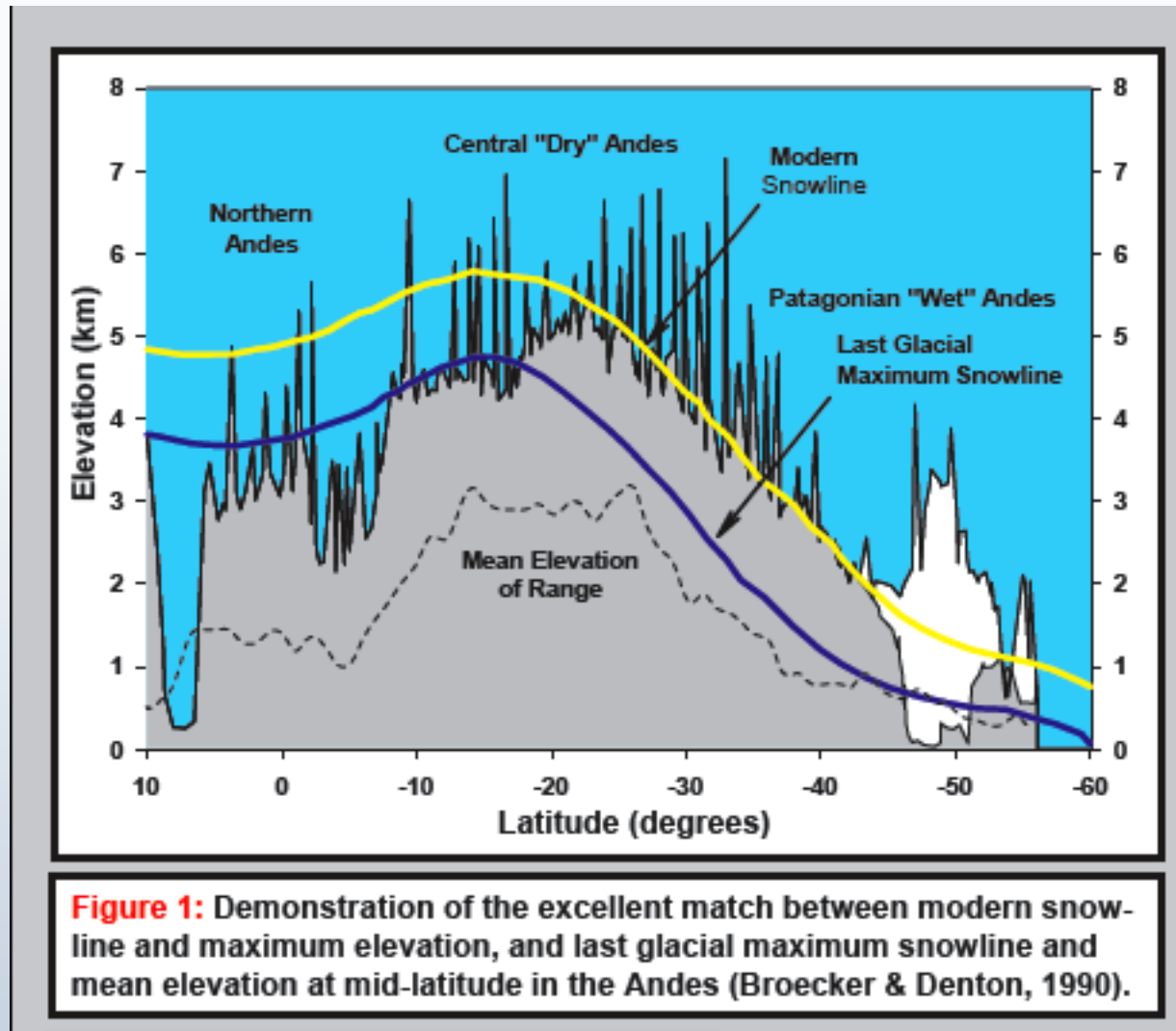


DJ Sauchyn, location unknown

- U-shaped glacial valley
- V-shaped fluvial valley

\* What is the footprint of the erosive processes?\*

# The "glacial buzzsaw"



- How fast is glacial erosion?

A photograph of a snowy mountain landscape. In the foreground, a white and yellow helicopter is landing on a snow-covered slope. To the right, a small metal cabin is visible. In the background, a large mountain peak is partially covered in snow. The sky is overcast.

**Thank You!**

**Questions?**

**EXTRA SLIDES**

# Varieties of glaciers



Upper Fremont Glacier, Rita Donham 2006  
southern Axel Heiberg Island, Canada, J. Alean, 1977

Grosser Aletsch glacier, Switzerland, Wikimedia commons

Antarctica satellite composite, Wikimedia commons

Icesheet

Icecap

Valley glacier

Piedmont glacier

Hanging glacier

Cirque glacier

- ELA elevation and topography determine accumulation area. This in part determines glacier size and morphology